

ATMOSPHERE DEPENDENCE OF THE THERMAL DECOMPOSITION OF MANGANITE, γ -MnO.OH

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ABSTRACT

Thermal analysis using TG and simultaneous DTA-TG-MS has been carried out on a natural manganite in atmospheres of oxygen, air, carbon dioxide, nitrogen and argon. The TG curve shows three distinct weight losses at approximately 300-400°, 500-600° and 800-900°C. The ratio of the first to second weight loss varies dramatically depending on atmosphere and, for oxidizing atmospheres, the heating rate in the range 2-50°C/min. In oxygen at 2°C/min, MnO₂ is the only product of the first (dehydroxylation) reaction but, with increasing heating rate and decreasing p O₂, Mn₂O₈ becomes a prominent co-product. In inert atmospheres, Mn₃O₈, Mn₂O₃ and Mn₃O₄ may all appear as products of the first reaction. Mn₃O₈ and MnO₂ decompose to Mn₂O₃ between 500° and 600°C and Mn₂O₃ decomposes to Mn₃O₄ at 800-900°C. Mn₃O₄ undergoes a β - γ transition at 1169°C and decomposes to MnO at 1450°C.

MATERIALS AND METHODS

The manganite used was a pure, crystalline variety from Laverock Braes, Aberdeenshire, which gave an X-ray pattern virtually identical with that of JCPDS card 8-99. This was ground to a particle size of <75 μ m and thermogravimetric analyses carried out on 10 mg samples in a Stanton Redcroft TG 770 microthermobalance in static air and in flowing oxygen, carbon dioxide and nitrogen. Heating rates were in the range 2° to 50°C/min. Separate TG runs were made to isolate products after significant weight losses for subsequent identification by X-ray powder photography (XPP). Simultaneous TG-DTA-MS at 15°C/min was carried out in argon on 10 mg samples using equipment described by Charsley et al. [1].

RESULTS AND REACTIONS

TG curves of manganite at 15°C/min in different atmospheres are given in Fig. 1. Three weight losses are shown in all cases, termed A (300-400°C), B (500-600°C) and C (800-900°C), but the ratio of A to B increases in the order oxygen < air < nitrogen < carbon dioxide. Table 1 summarizes results from TG runs carried out under atmospheres ranging from "maximum oxidation" (slow heating under flowing oxygen) to neutral (flowing argon). The Table demonstrates well the systematic increase of weight loss A at the expense of B with decreasing p O₂. Weight loss C increases to a maximum of 3.2% in static air at 15°C/min and then decreases to 2.3% in argon, while the total

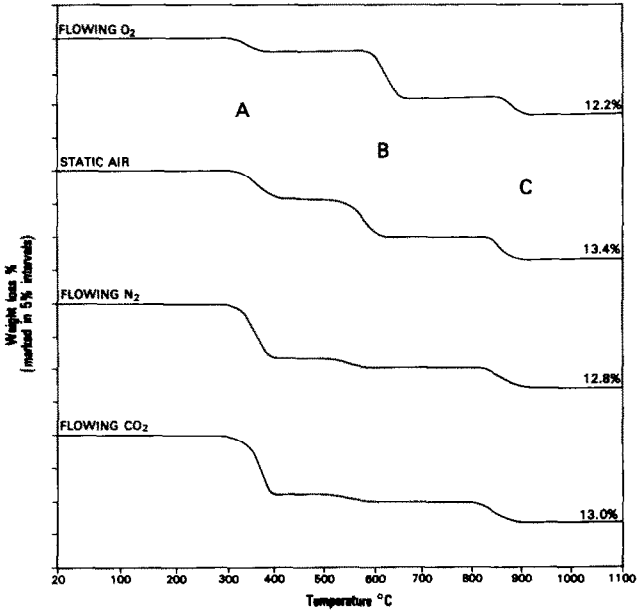


Fig. 1. TG curves of manganite under different flowing gas atmospheres at 15°C/min.

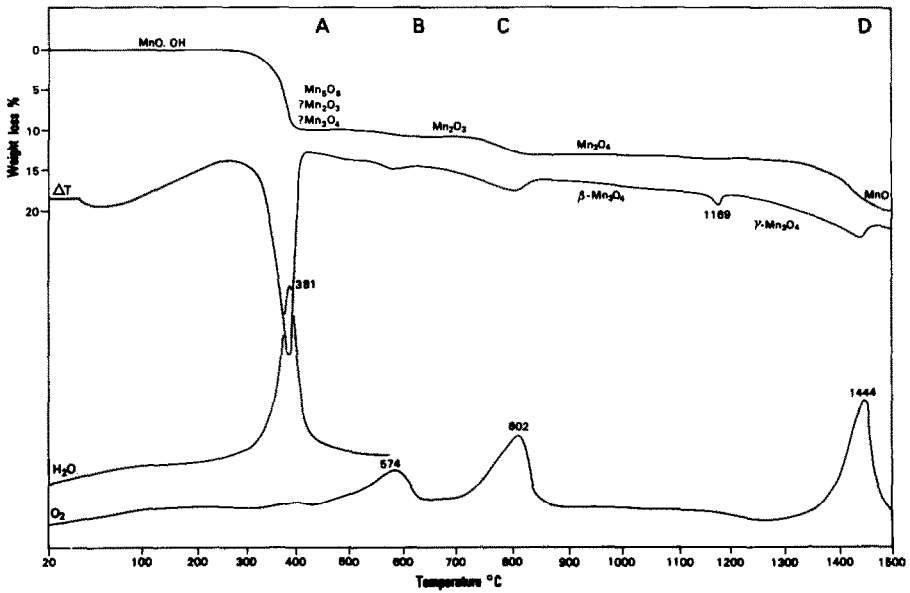
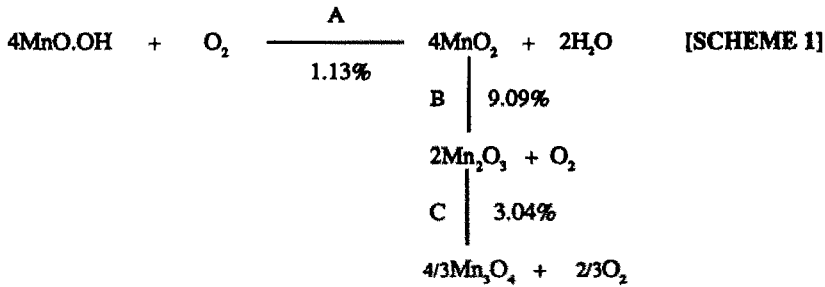
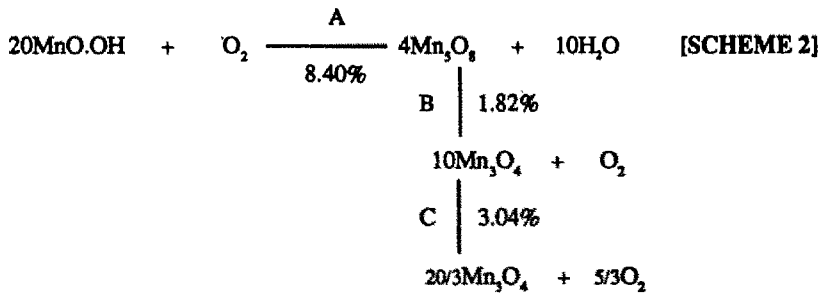


Fig. 2. Simultaneous DTA-TG-MS curves of manganite in flowing argon at 15°C/min.

weight loss to 1000°C generally increases with decrease in p_{O_2} . Weight losses from slow heating rates in oxygen are in general agreement with the reaction scheme:



With decreasing p_{O_2} , Mn_3O_4 was identified by XPP as an increasingly important co-product of weight loss A. A possible scheme to account for the formation of this phase, and subsequent reactions, is:

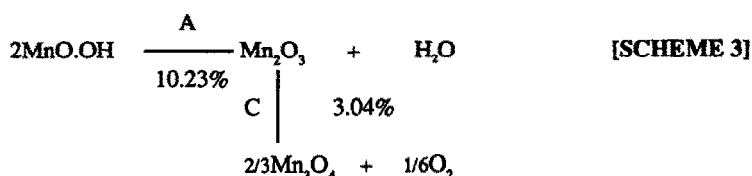


Atmosphere	Heating rate (°C/min)	Weight loss (%)			Total
		A 300-400°C	B 500-600°C	C 800-900°C	
Oxygen	2	1.2	8.2		
	5	1.5	7.9	2.1	11.5
	15*	2.3	7.3	2.6	12.2
	50	2.95	6.15		
Static air	5	3.0	7.0		
	15*	4.4	5.8	3.2	13.4
	50	5.5	3.8	3.1	13.4
Nitrogen	15*	8.5	1.45	2.85	12.8
Carbon dioxide	15*	9.0	1.3	2.7	13.0
Argon	15†	10.0	0.9	2.3	13.2

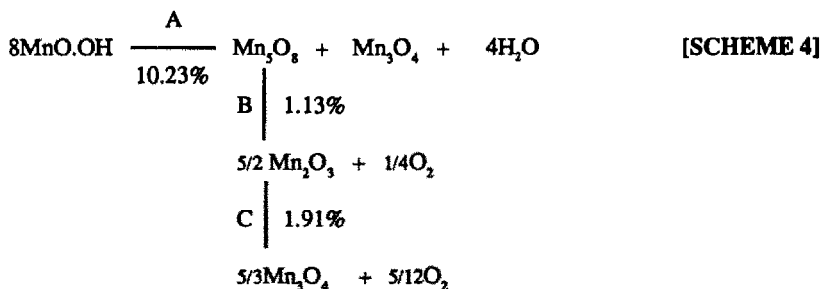
Table 1. Weight losses shown by manganite under different flowing gas atmospheres. †TG curves shown in Fig. 1; *TG curve shown in Fig. 2.

Schemes 1 and 2 probably occur simultaneously in the presence of oxygen, with scheme 2 predominating with decreasing p_{O_2} . As both MnO_2 and Mn_2O_3 decompose to Mn_2O_3 between 500° and 600°C [2,3], it is possible to estimate the proportions of the two phases formed during weight loss A from its actual value - for instance, the amount of Mn_2O_3 formed in static air varied from 25% at 5°C/min to 60% at 50°C/min.

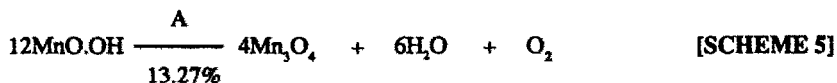
In carbon dioxide, nitrogen and argon, however, weight loss A exceeds 8.40% so a different reaction scheme must operate. Oswald et al. [3] found that manganite transformed topotactically to Mn_2O_3 in nitrogen, but this is not possible stoichiometrically unless another Mn-oxide of lower valency accompanies the Mn_2O_3 . XPP examination of the product from weight loss A in nitrogen identified both Mn_2O_3 and Mn_3O_4 . The formation of Mn_2O_3 alone as the product of weight loss A could be explained by reaction scheme 3:



Although this would account for the larger weight loss A shown in inert atmospheres, there would be no weight loss B as the Mn_2O_3 formed decomposes to Mn_3O_4 between 800° and 900°C. Table 1 shows that there is a weight loss B in inert atmospheres, this decreasing from nitrogen to argon. Lee et al. [4] demonstrated that, under vacuum, $MnO.OH$ decomposed to equimolar amounts of Mn_2O_3 and Mn_3O_4 , which forms the basis for reaction scheme 4 below:



Weight losses A, B and C are close to those shown under argon in Table 1, as is the overall loss to 1000°C of 13.27%. Simultaneous DTA-TG-MS curves in argon (Fig. 2) show water evolution during weight loss A and oxygen evolution during weight losses B and C, in accordance with scheme 4. There also appears to be a slight oxygen evolution during weight loss A, which could be accounted for by an accompanying reaction of the type:



The generation of oxygen from such reactions would explain the discrepancies between weight losses in the different "inert" atmospheres in Table 1.

Fig. 2 also shows two further high-temperature reactions of manganite - the β - γ transition of the Mn_3O_4 reaction product at 1169°C [2] and decomposition of this to MnO above 1400°C.

CONCLUSIONS

Reaction schemes have been proposed that explain most of the variations shown by the TG curve of manganite in different atmospheres. Topotactic decomposition to Mn_3O_4 occurs at low p_{O_2} between 300-400°C [3], but other Mn-oxides appear as products of the initial decomposition at high p_{O_2} and in inert atmospheres and these influence subsequent reactions. Variations in TG curves found at different heating rates in the same atmosphere are due to differences in rate of expulsion of the product gases (O_2 and H_2O) vs. rate of inward diffusion of atmosphere gas. Discrepancies in Table 1 between weight losses given in different "inert" atmospheres may be due, at least in part, to differences in diffusivity of the gases used, and thus variations in the self-generated sample atmosphere. This demonstrated dependence of decomposition mechanism on atmosphere and heating rate has important implications to the quantitative determination of manganite or similar artificial analogues by thermal methods, particularly TG.

ACKNOWLEDGMENTS

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